

Structural constraints on the evolution of the south-eastern Mwanesi Greenstone Belt and adjacent granitoids, central Zimbabwe Craton: implications for gold mineralisation

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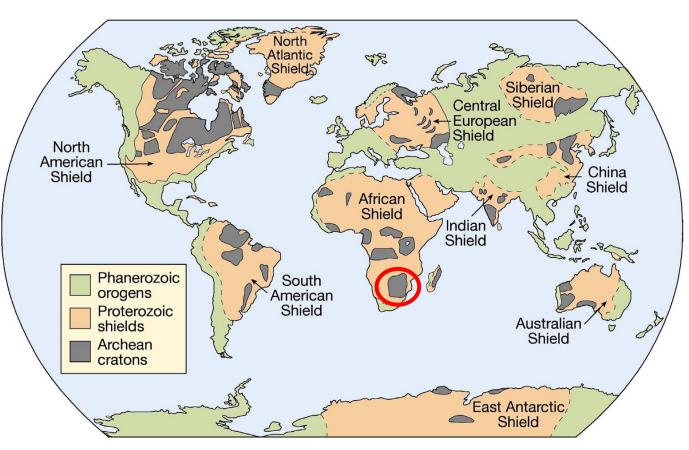
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IRP BuCoMO



Archaean greenstone belts



Furnes *et al.* (2013)

- Common in Archaean cratons
- Tectonic evolution still debatable (e.g. Cawood et al.

2018; Gapais, 2018; Brown et al. 2020)

- Host orogenic gold deposits (Groves and Foster, 1991)
- Mineralising fluids source still controversial (Goldfarb)

and Groves, 2015; Groves et al. 2019a, 2019b)

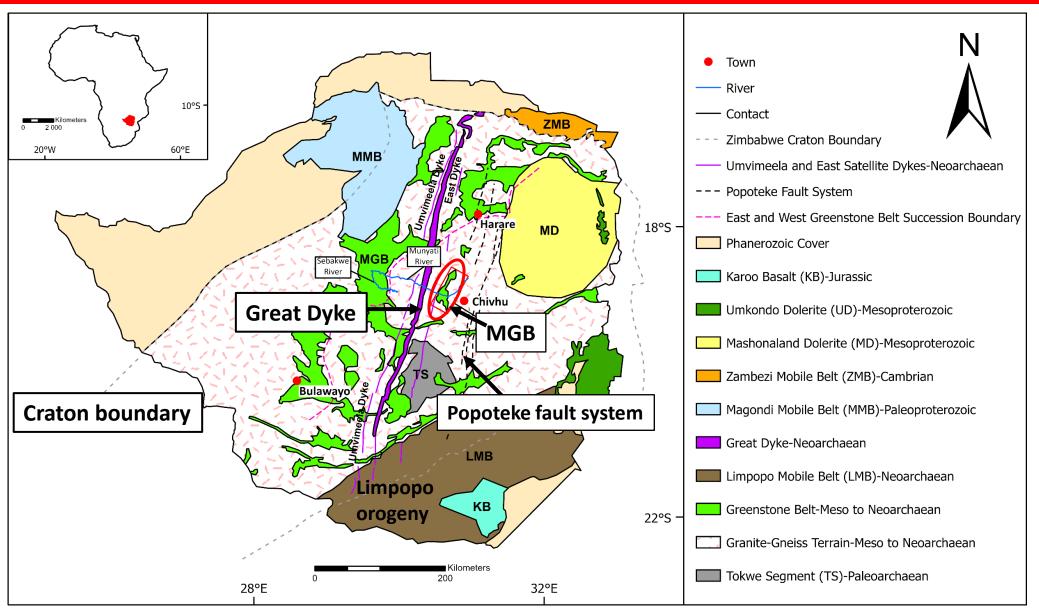
Aims of this study

- Deformation of the Mwanesi Greenstone Belt (MGB) and the surrounding granitoids
- 2. The link between deformation and gold mineralisation (controls and timing)
- 3. Suggest a unified tectonic model for the evolution of the MGB

Why the MGB?

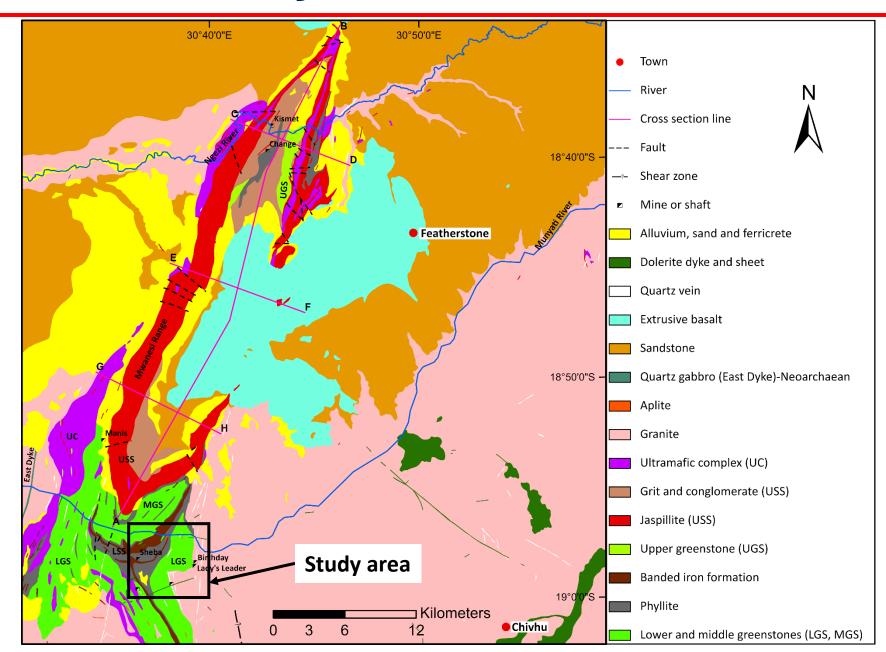
- One of the least studied greenstone belts in the Zimbabwe Craton (ZC)
- Its low gold endowment compared to other greenstone belts of the ZC

Geology of the Zimbabwe Craton



Modified after Markwitz et al. (2010), Western and eastern succession greenstone belts division is after Wilson (1979).

Geology of the Mwanesi Greenstone Belt



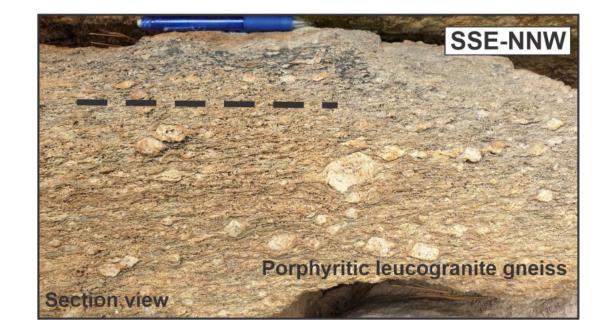
Modified after Worst (1962)

Lithostratigraphy – Adjacent granitoids



- Gneissic granites and minor isotropic granite
- Medium to coarse-grained, porphyritic, largely leucocratic
- Grey gneiss intruded by leucogranite gneisses
- Granite gneiss bands, and crosscutting dikes





Lithostratigraphy - Lower Greenstone Series

- **Metasedimentary rocks** pelitic schists
- Volcanic rocks pillow and amygdaloidal basalts, pyroclasticbreccia, felsic volcanic rocks
- Intrusive rocks mafic rocks, and dikes

Metasedimentary rocks





Volcanic rocks



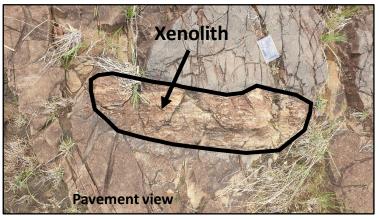






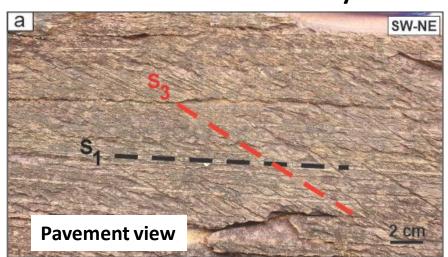
Intrusive rocks

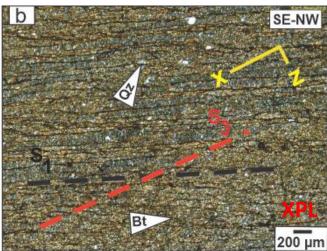




Lithostratigraphy - Lower Sedimentary Series

Phyllite

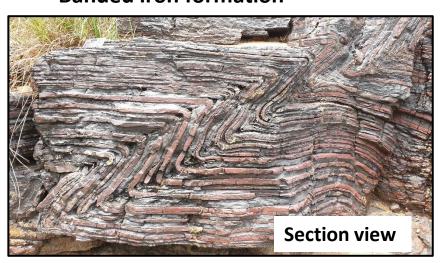




- Phyllite and BIF
- BIF intruded by subvolcanic basalts
- Intrusive greenstones (Worst, 1962)

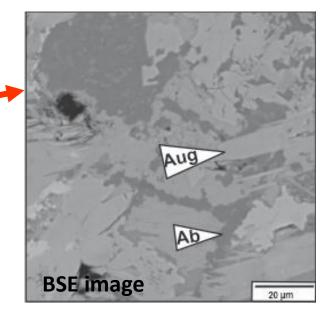
Subvolcanic sill

Banded iron formation



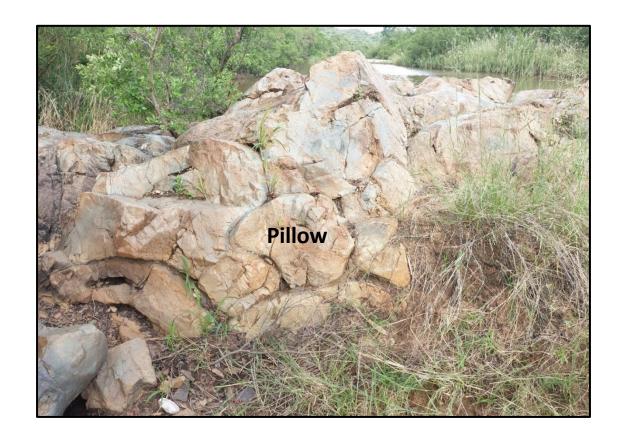
Banded iron formation

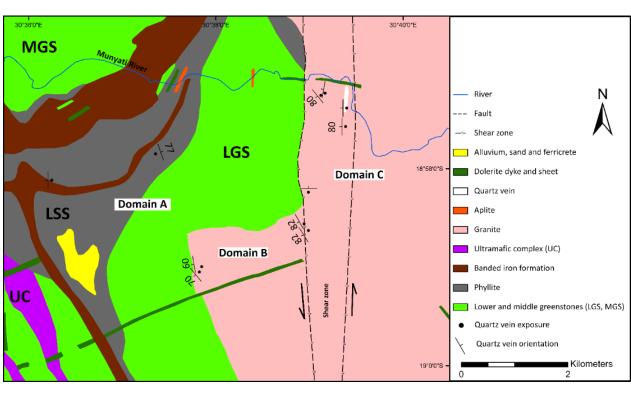




Lithostratigraphy - Middle greenstone Series

- Characterised by basaltic rocks, locally pillowed
- MGS and LGS pillow basalts are lithologically similar
- Minor textural differences





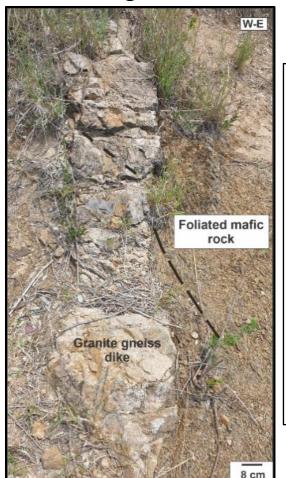
Modified after Worst (1962)

Granite-greenstone age relationship

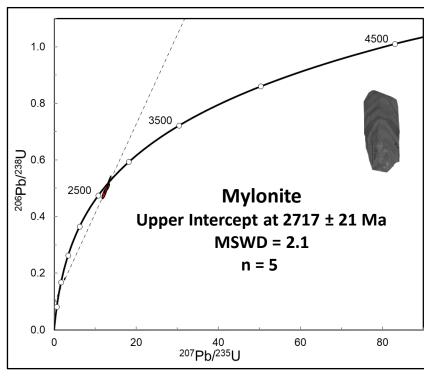
- All granitoids younger than the MGB Worst (1962)
- Not all granitoids are younger (this study)

MGS LGS Alluvium, sand and ferricrete Dolerite dyke and sheet Domain C Domain A Ultramafic complex (UC) Domain B wer and middle greenstones (LGS, MGS) Quartz vein exposure **Mhou Shear Zone (MSZ)**

Granite gneiss dike

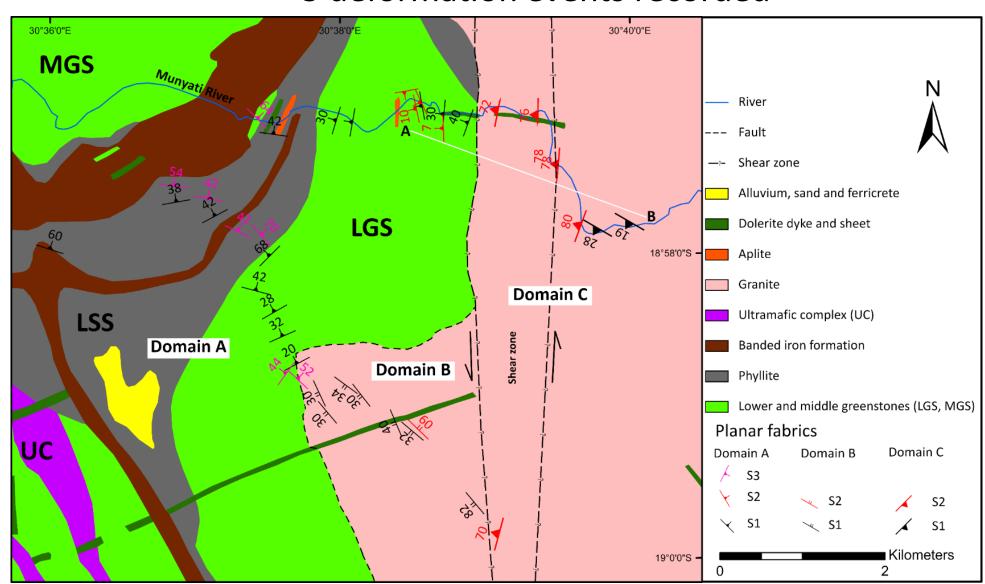


Mylonite sample



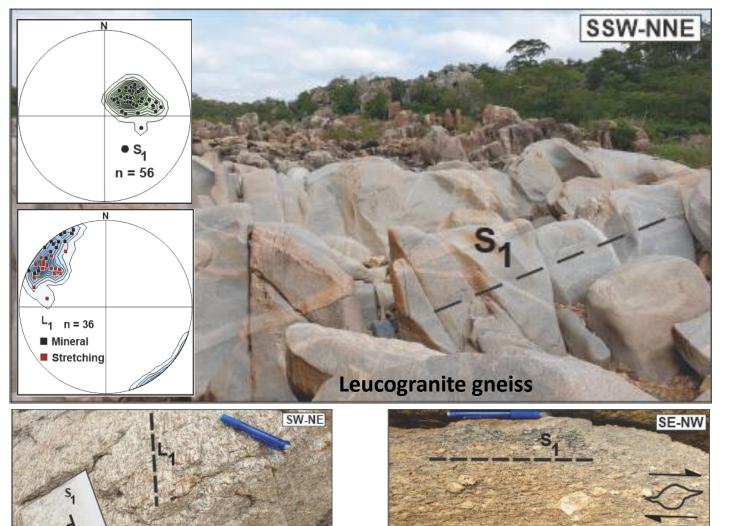
Deformation

• 3 deformation events recorded

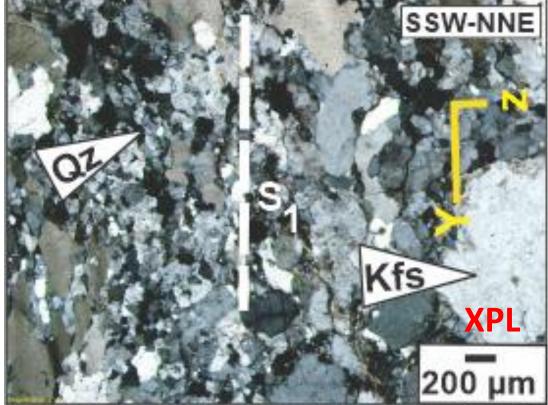


Modified after Worst (1962)

D₁ event – granite gneisses



Ductile deformation (300-450°C)

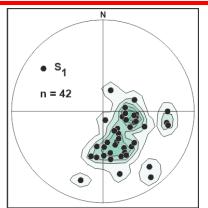


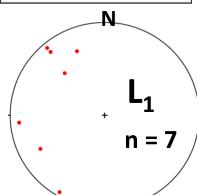
Porphyritic leucogranite gneiss

Section view

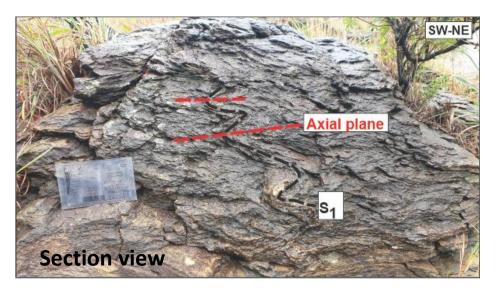
Section view

D₁ event – supracrustal rocks

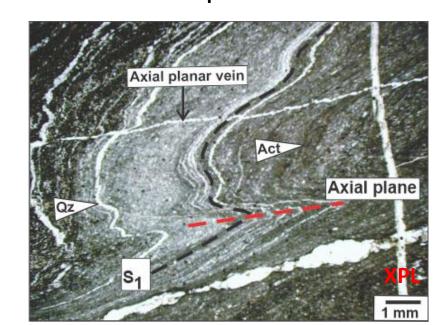


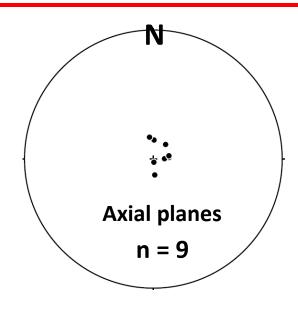


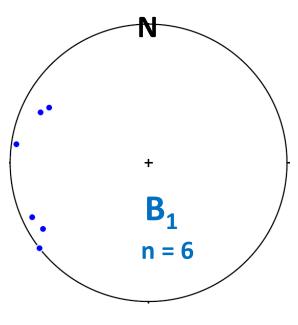




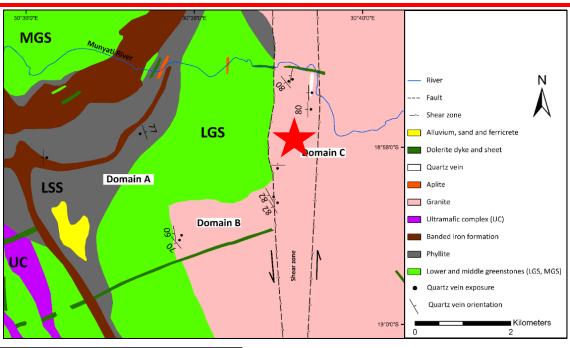
Medium-laminated pelitic schist







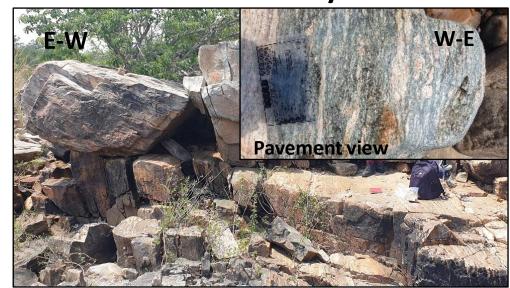
D₂ event - Mhou Shear Zone (MSZ)



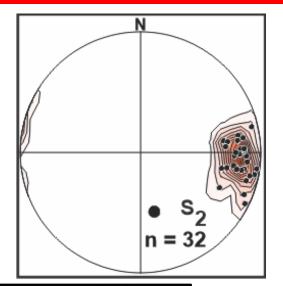
Ultramylonite

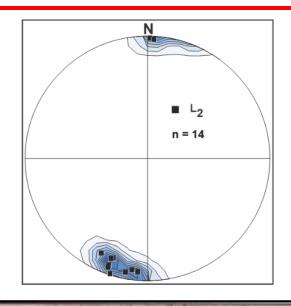


Protomylonite

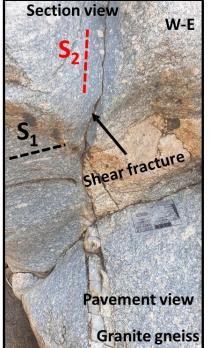


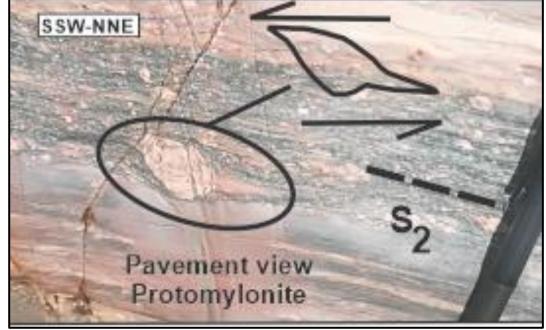
D₂ event

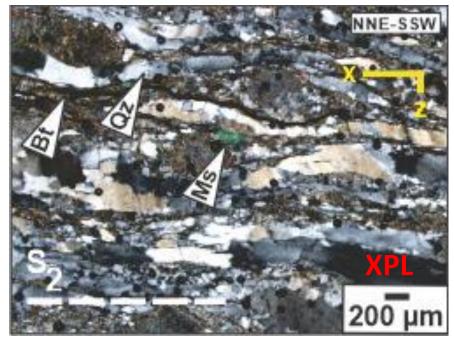




- Crystal plastic deformation in quartz, quartz ribbons
- Ductile deformation (400-550 °C)
- Between NW-SE and NNW-SSE directed shortening



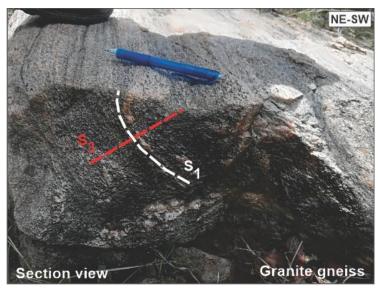


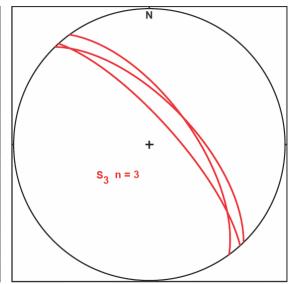


D₃ event

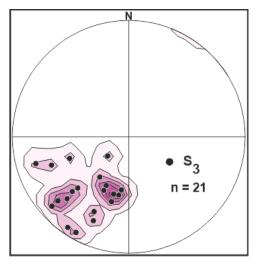
- Cm-dm-scale box, open and inclined **F**₃ folds
- Moderately NE-dipping axial planes and axial planar cleavage S₃
- NE-SW shortening

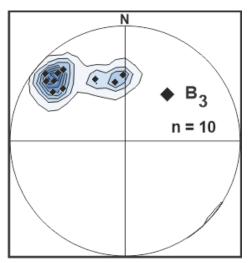
Granite gneisses W of the MSZ

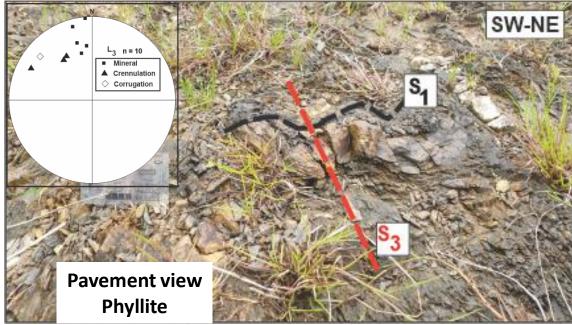




Supracrustal rocks





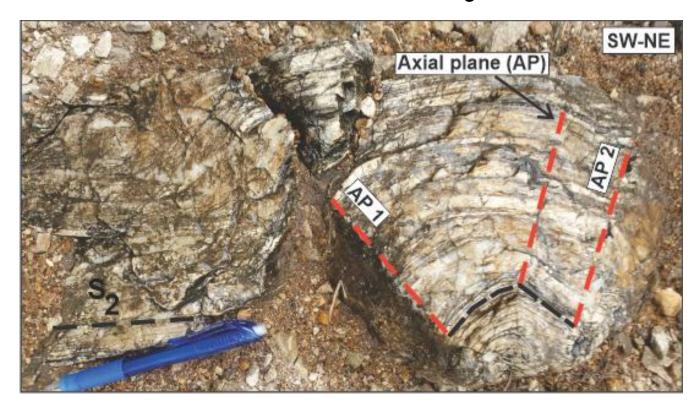


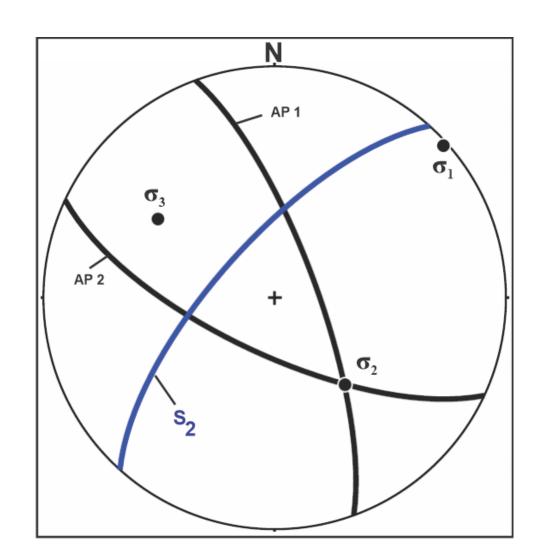
D₃ event

Paleostress direction reconstruction after Ramsay and Huber

(1987) – box folds

• NE-SW shortening inferred from **F**₃ box folds



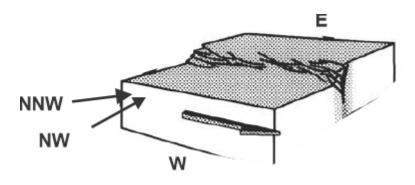


Interpretation of deformation events

- D_{I}
- Contractional regime with a top to the NW simple shear component
- Ductile deformation

 D_2

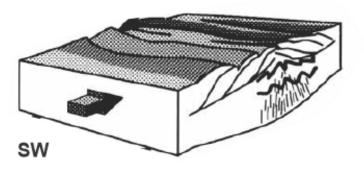
Wrench dominated transpression



- Medium grade mylonites
- Ductile deformation

D₃
Contractional regime

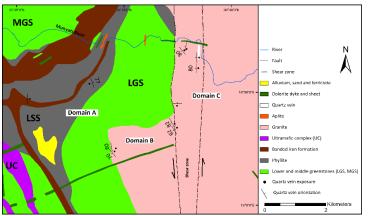
ΝE



Brittle-ductile deformation

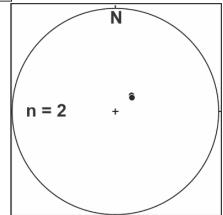
Sketches modified after Fossen et al (1994)

Quartz veins

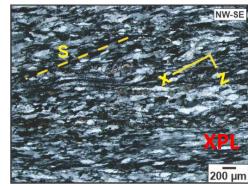


Veins W of the MSZ

- M-thick, shallowly SW-dipping
- Pre- or early syn-D₁

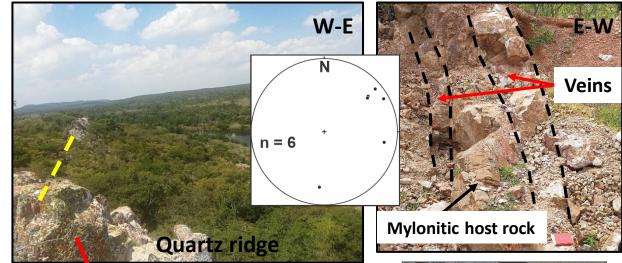




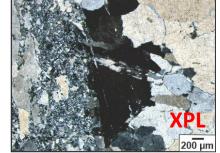


Veins in the MSZ

- Steeply SW-W- and N dipping
- Dm-thick; m-thick up to ~40 m thick
- Post-D₂ veins
- More than one fluid injection event

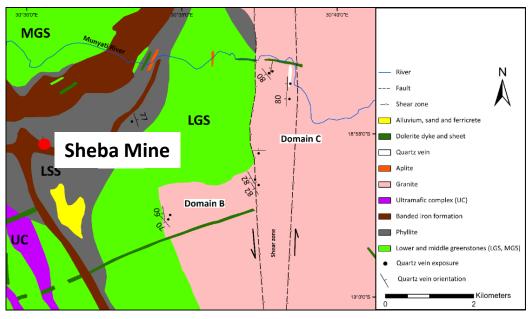






Gold mineralisation

- Gold mineralisation occurs in both the MGB and adjacent granitoids
- Gold mineralisation hosted by quartz veins
- Not enough information about mineralisation in the MGB
- In the MSZ, post D₂ veins mineralized (this study)
- Mineralised veins generally SW-dipping (this study)
- Very low grades, averaging 5g/t (From small scale miners)



Au assay results

MGB – Sheba Mine Veins W of MSZ

3 g/t 0.01 g/t

MSZ veins

W-dipping veins

Average – 0.01 g/t

SW-dipping veins

0.03 g/t & 0.13 g/t

Take home messages

- Polyphase deformation in the MGB 3 deformation events lower-upper greenschist facies
- $\mathbf{D_1}$ recumbent folding in the supracrustal rocks and top to the NW-shearing in the underlying granite gneisses, $\mathbf{D_2}$ wrench dominated transpression and $\mathbf{D_3}$ contractional tectonic regime
- Two generations of quartz veins recognized i.e. pre- or early syn-D₁ and post D₂
- Gold mineralisation hosted in post **D**₂ quartz veins, grades generally low

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Thank you



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